

## CHAPTER FOUR

### GRADUALLY VARIED FLOW (GVF)

#### 4.1 General

A steady non-uniform flow in a prismatic channel with gradual changes in its water-surface elevation is termed as gradually-varied flow (GVF). The back water produced by a dam or weir across a river and the drawdown produced at a sudden drop in a channel are few typical examples of GVF. In a GVF, the velocity varies along the channel and consequently the bed slope, water surface slope, and energy slope will all differ from each other. Regions of high curvature are excluded in the analysis of this flow.

The two basic assumptions involved in the analysis of GVF are:

1. The pressure distribution at any section is assumed to be hydrostatic. This follows from the definition of the flow to have a gradually-varied flow surface. As gradual changes in the surface of the curvature give rise to negligible normal acceleration, the departure from the hydrostatic pressure distribution is negligible.

Slope of the channel is small.

In GVF bed slope, water surface slope and energy slope are all different. Hydrostatic normal acceleration = 0

2. The resistance to flow at any depth is assumed to be given by the corresponding uniform flow equation, such as Manning's formula, with the condition that the slope term to be used in the equation is the energy slope and not the bed slope. Thus, if in a GVF the depth of flow at a section is  $y$ , the energy slope  $S_f$  is given by:

$$S_f = \frac{n^2 V^2}{R^{4/3}}$$

According to the assumption: the Manning's formula can be used to evaluate the energy slope of GVF

#### 4.2 Differential equation of GVF

Consider the total and specific energy  $H$  and  $E$  respectively of a gradually-varied flow in a Channel of small slope and  $\alpha=1.0$  in elementary length  $dx$  of the channel:-

$$H = Z + E = Z + y + \frac{V^2}{2g} \quad [1]$$

A schematic sketch of a gradually-varied flow is shown in figure 5.1. Since the water surface, in General, varies in the longitudinal ( $x$ ) direction, the depth of flow and total energy are functions of  $x$ . Differentiation the above equation with respect to  $x$  shows that:-

$$\frac{dH}{dx} = \frac{dZ}{dx} + \frac{dE}{dx} = \frac{dZ}{dx} + \frac{dy}{dx} + \left( \frac{V^2}{2g} \right) \frac{d}{dx} \quad [2]$$

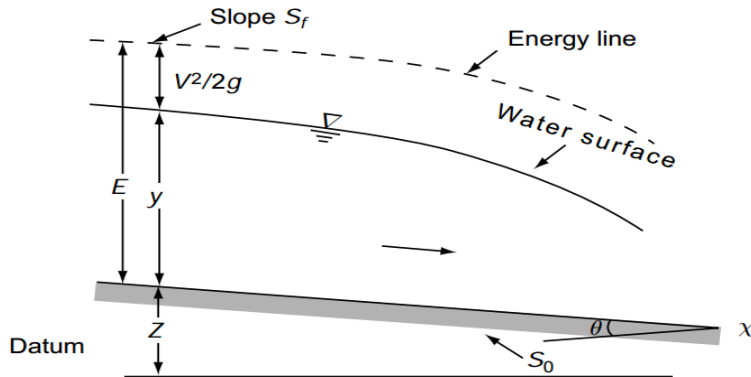


Figure 4.1 Schematic sketch of GVF

In this equation the meaning of each term is as follows:-

1. — Represents the energy slope. Since the total energy of the flow always decreases in the direction of the motion, it is common to consider the slope of the decreasing energy line as positive. Denoting it by  $S_f$ :-

$$\frac{dH}{dx} = -S_f$$

2. — Denotes the bottom slope. It is common to consider the channel slope with bed elevations decreasing in the downstream direction as positive. Denoting it as  $S_o$ .

$$\frac{dZ}{dx} = -S_o$$

3. — Represents the water surface slope relative to the bottom of the channel.

$$\begin{aligned}
 4. \quad \left( \frac{V^2}{2g} \right) \frac{d}{dx} &= \left( \frac{Q^2}{2gA^2} \right) \frac{d}{dy} \frac{dy}{dx} \\
 &= - \frac{Q^2}{gA^3} \frac{dA}{dy} \frac{dy}{dx} \quad \text{Since } dA/dy=T \\
 \left( \frac{V^2}{2g} \right) \frac{d}{dx} &= - \frac{Q^2 T}{gA^3} \frac{dy}{dx} \quad \text{Equation [2] can be written as:} \\
 -S_f &= -S_o + \frac{dy}{dx} - \frac{Q^2 T}{gA^3} \frac{dy}{dx} \quad \text{Rearranging} \\
 \frac{dy}{dx} &= \frac{S_o - S_f}{1 - \frac{Q^2 T}{gA^3}} \quad [3]
 \end{aligned}$$

This forms the basic differential equation of GVF and is also known as the dynamic equation of GVF. If a value of the kinetic energy correction factor  $\alpha$  greater than unity is to be used,

$$\frac{dy}{dx} = \frac{S_o - S_f}{1 - \frac{\alpha Q^2 T}{gA^3}} = \frac{S_o - S_f}{1 - \alpha F^2}$$

## Other forms of GVF equation

a) If  $K$  = conveyance at any depth  $y$  and  $K_o$  = conveyance corresponding to the normal depth  $y_o$ , then

$$K = \frac{Q}{\sqrt{S_f}} \text{ by assumption 2 of GVF}$$

$$\text{and } K_o = \frac{Q}{\sqrt{S_o}} \text{ uniform flow formula}$$

$$\frac{S_f}{S_o} = \frac{K_o^2}{K^2}$$

Similarly if  $Z$  = section factor at a depth  $y$  and  $Z_c$  = section factor at the critical depth  $y_c$ ,

$$Z^2 = \frac{A^3}{T} \text{ and } Z_c^2 = \frac{A_c^3}{T_c} = \frac{Q^2}{g}$$

$$\text{Hence } \frac{Q^2 T}{g A^3} = \frac{Z_c^2}{Z^2}$$

Equation 3, the differential equation of GVF can be written as:

$$\frac{dy}{dx} = S_o \frac{1 - \frac{S_f}{S_o}}{1 - \frac{Q^2 T}{g A^3}}$$

$$\frac{dy}{dx} = S_o \frac{1 - \left(\frac{K_o}{K}\right)^2}{1 - \left(\frac{Z_c}{Z}\right)^2} \text{ usefull in developing direct integration technique}$$

b.) if  $Q_n$  represents the normal discharge at a depth  $y$  and  $Q_c$  denotes the critical discharge at the same depth  $y_c$  and  $Q$  is the given discharge of the GVF at a depth  $y$ .

$$Q_n = K\sqrt{S_o}, \quad Q = K\sqrt{S} \text{ and}$$

$$Q_c = Z\sqrt{g} \text{ using those equations,}$$

$$\frac{dy}{dx} = S_o \frac{1 - (Q/Q_n)^2}{1 - (Q/Q_c)^2}$$

c.) total energy

$$H = E + Z$$

$$\frac{dH}{dx} = \frac{dE}{dx} + \frac{dZ}{dx} \quad -S_f = -S_o + \frac{dE}{dx}$$

$$\frac{dE}{dx} = S_o - S_f$$

$$\Delta X = \frac{\Delta E}{S_o - S_f} \text{ usefull in developing numerical technique for GVF profile computations.}$$

### 4.3 Classification of flow profiles

In a given channel,  $y_0$  and  $y_c$  are two fixed depths if  $Q$ ,  $n$  and  $S_0$  are fixed. Also there are three possible relations between  $y_0$  and  $y_c$  as:-

(i)  $y_0 > y_c$ , (ii)  $y_0 < y_c$  and (iii)  $y_0 = y_c$ . further there are two cases where  $y_0$  doesn't exist, i.e when a) the channel bed is horizontal ( $S_0=0$ ), b) when the channel has an adverse slope, ( $S_0$  is -ve). Based on the above, the channels are classified into five categories as indicated in table 4.1.

Table 4.1 classification of channels

No	Channel category	Symbol	Characteristic condition	Remark
1	Mild slope	M	$y_0 > y_c$	Subcritical flow at normal depth
2	Steep slope	S	$y_0 < y_c$	Supercritical flow at normal depth
3	Critical slope	C	$y_0 = y_c$	Critical flow at normal depth
4	Horizontal bed	H	$S_0=0$	Cannot sustain uniform flow
5	Adverse slope	A	$S_0 < 0$	Cannot sustain uniform flow

For each of **five** categories of channels, lines representing the critical depth and normal depth (if it exists) can be drawn in the longitudinal section. Those would divide the whole flow space into three regions as:-

**Region 1: space above the top most line**

**Region 2: space between top line and the next lower line**

**Region 3: space between the second line and the bed**

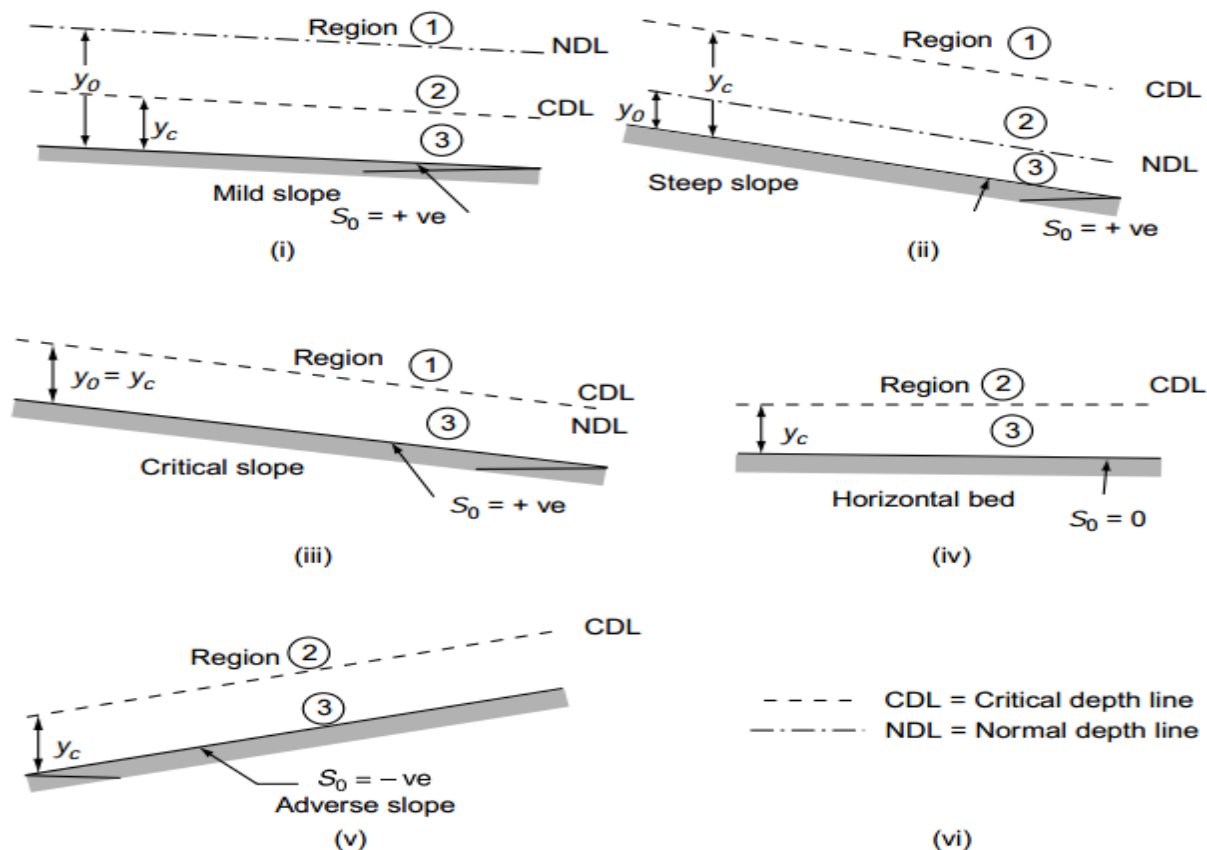


Figure 4.2 regions of flow profiles

Depending upon the channel category and region of flow, the water-surface profiles will have either of the following characteristic shapes.

1. **Back water curves**: if the depth of flow increases in the direction of flow.

2. **Drawdown curves**: if the depth of flow decreases in the direction of flow.

The dynamic equation of GVF expresses the longitudinal surface slope of flow with respect to the channel bottom is given by:-

$$\frac{dy}{dx} = S_o \frac{1 - \left(\frac{K_o}{K}\right)^2}{1 - \left(\frac{Z_c}{Z}\right)^2}$$

**1. Back water curves**  $dy/dx$  is positive

**Case 1** if the numerator  $> 0$  and denominator  $> 0$

$$1 - \left(\frac{K_o}{K}\right)^2 > 0 \text{ and } 1 - \left(\frac{Z_c}{Z}\right)^2 > 0$$

That is  $K > K_o$  and  $Z > Z_c$

**Case 2** if the numerator  $< 0$  and denominator  $< 0$

$$1 - \left(\frac{K_o}{K}\right)^2 < 0 \text{ and } 1 - \left(\frac{Z_c}{Z}\right)^2 < 0$$

That is  $K < K_o$  and  $Z < Z_c$

For channel of the first kind,  $K$  is a single valued function of  $y$ , and hence

$$\frac{dy}{dx} > 0 \text{ if } y > y_o \text{ and } y > y_c \text{ or}$$

$$y < y_o \text{ and } y < y_c$$

Three cases:  $y > y_o > y_c$   
 $y > y_c > y_o$   
 $y > y_c = y_o$

Three cases:  $y_c > y_o > y$   
 $y_o > y_c > y$   
 $y_c = y_o > y$

**2. Drawdown curves**  $dy/dx$  is negative

**Case 1**

$$1 - \left(\frac{K_o}{K}\right)^2 > 0 \text{ and } 1 - \left(\frac{Z_c}{Z}\right)^2 < 0$$

That is  $K > K_o$  and  $Z < Z_c$

**Case 2**

$$1 - \left(\frac{K_o}{K}\right)^2 < 0 \text{ and } 1 - \left(\frac{Z_c}{Z}\right)^2 > 0$$

That is  $K < K_o$  and  $Z > Z_c$

For channel of the first kind,  $K$  is a single valued function of  $y$ , and hence

$$\frac{dy}{dx} > 0 \text{ if } y_c > y > y_o \text{ or}$$

$$y_o > y > y_c$$

Two cases:  $y_c > y > y_o$   
 $y_o > y > y_c$

Further, to assist in determination of flow profiles in various regions, the behavior of  $dy/dx$  at certain key depths is noted by studying the differential equation of GVF.

$$\frac{dy}{dx} = S_o \frac{1 - \left(\frac{K_o}{K}\right)^2}{1 - \left(\frac{Z_c}{Z}\right)^2}$$

As  $y \rightarrow y_o$   $dy/dx \rightarrow 0$

The water surface meets the normal depth line asymptotically

As  $y \rightarrow \infty$   $dy/dx \rightarrow S_o$

The water surface meets a very large depth as a horizontal asymptote /tends to be horizontal/

As  $y \rightarrow y_c$   $dy/dx \rightarrow \infty$

The water surface meets the CDL vertically /normally/

In reality, high curvatures at the critical depth zone violate the assumption of GVF, then the profile has to end a short distance away from the  $y_c$  location. At critical depth the curves are indicated by dashed lines to remind that the GVF equation is strictly not applicable in that Neighborhood.

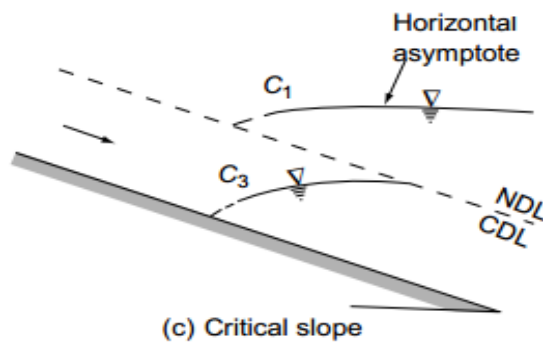
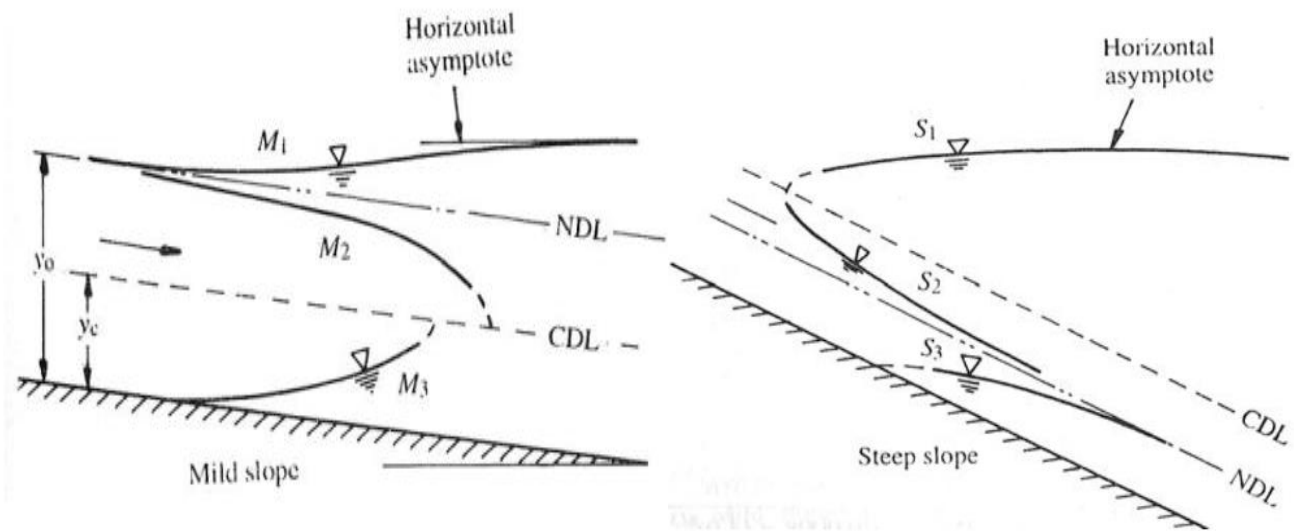
As  $y \rightarrow 0$   $dy/dx \rightarrow \infty$

The water surface meets the channel bottom normally.

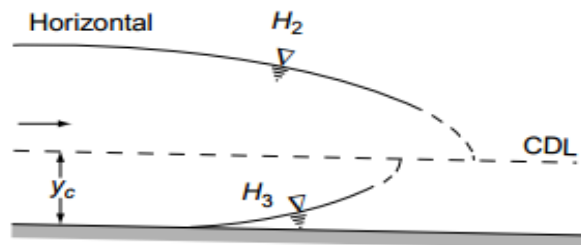
Based on the above information, the various possible gradually-varied flow profiles are grouped into *twelve types*.

No	Channel	Region	condition	Type
1	Mild slope	1	$y > y_o > y_c$	M <sub>1</sub>
		2	$y_o > y > y_c$	M <sub>2</sub>
		3	$y_o > y_c > y$	M <sub>3</sub>
2	Steep slope	1	$y > y_c < y_o$	S <sub>1</sub>
		2	$y_c > y > y_o$	S <sub>2</sub>
		3	$y_c > y_o > y$	S <sub>3</sub>
3	Critical slope	1	$y > y_o = y_c$	C <sub>1</sub>
		3	$y < y_o = y_c$	C <sub>3</sub>
4	Horizontal bed	2	$y > y_c$	H <sub>2</sub>
		3	$y < y_c$	H <sub>3</sub>
5	Adverse slope	2	$y > y_c$	A <sub>2</sub>
		3	$y < y_c$	A <sub>3</sub>

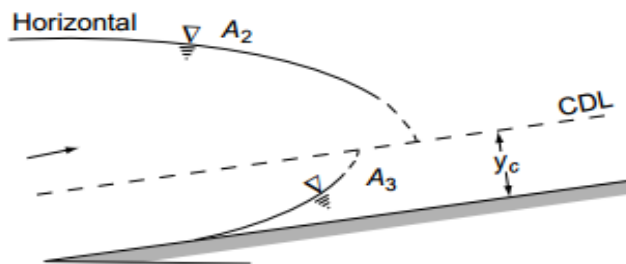
In reality the GVF profiles, especially M<sub>1</sub>, M<sub>2</sub> and H<sub>2</sub> profiles, are very flat.



(c) Critical slope



(d) Horizontal bed



(e) Adverse slope

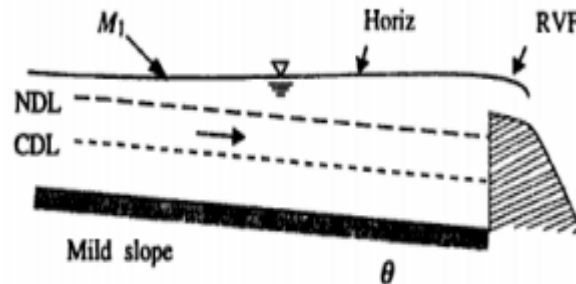
All curves in region 1 have **positive slopes**, known as **back water curves** and all curves in **region 2** have **negative slopes** and are referred as **draw down curves**.



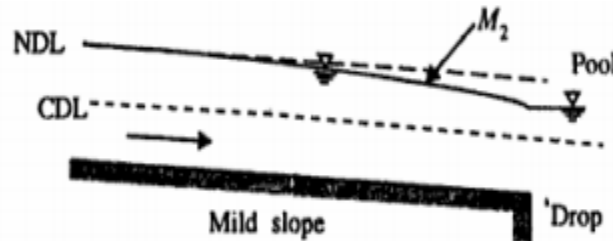
#### 4.4 same features of flow profiles

##### A). Type M flow profiles

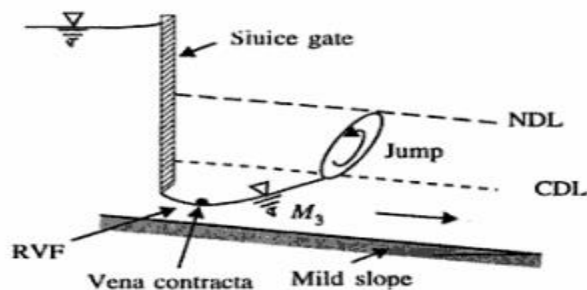
the most common of all GVF profiles is the M1 type, which is subcritical flow condition. Obstruction to the flow, such as weirs, dams, control structures and natural features, such as bends, produce M1 curves. These extend to several kilometers upstream merging with the normal depth.



The M2 profiles occur at a sudden drop in the bed of the channel, at constriction type of transitions and at the canal outlet into pools.

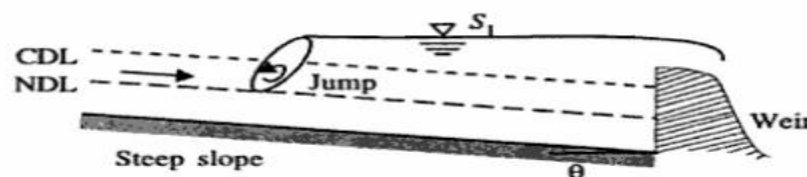


Where a supercritical stream enters a mild-slope channel, the M3 type of profile occurs. The flow leading from a spillway or a sluice gate to a mild slope forms a typical example. The beginning of the M3 curve is usually followed by a small stretch of rapidly varied flow and the down-stream is generally terminated by a hydraulic jump. Compared to M1 and M2 profiles, M3 curves are of relatively short length.



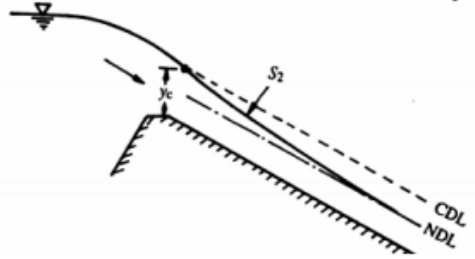
##### B). Type S profiles

The S1 profile is produced when the flow from a steep channel is terminated by a deep pool created by an obstruction, such as weir or dam. At the beginning of the curve, the flow changes from the normal depth (super critical flow) to subcritical flow through a hydraulic jump. The profiles extend downstream with a positive water-surface slope to reach a horizontal asymptote at the pool elevation.

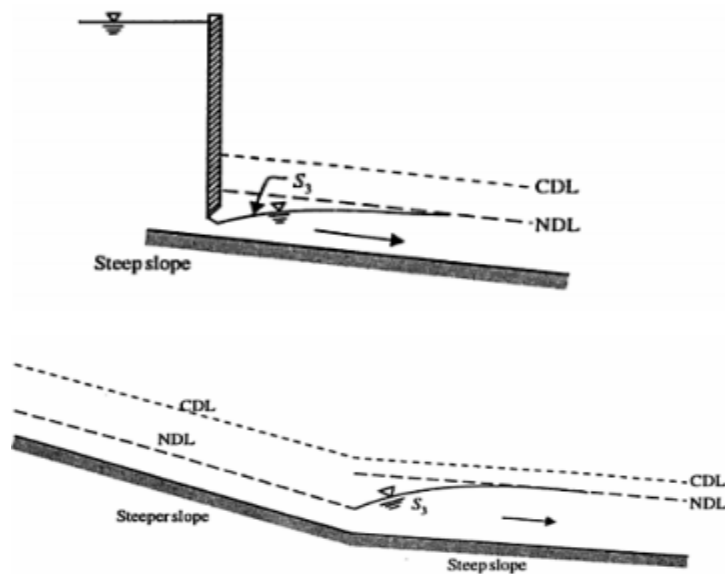




Profile of S2 type occurs at the entrance region of a steep channel leading from a reservoir and at the break of grade from mild slopes to steep slope. Generally S2 profiles are of short distance.



Free from a sluice gate with a steep slope on its downstream is of the S3 type. The S3 curve also results when a flow exists from a steeper slope to a less steep slope.

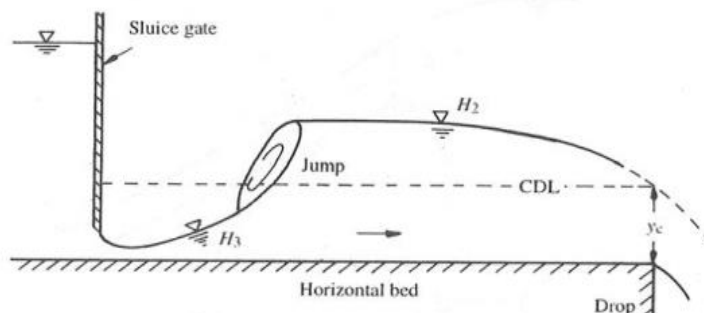


### c). Type C profiles

C1 and C3 curves are very rare and are highly unstable

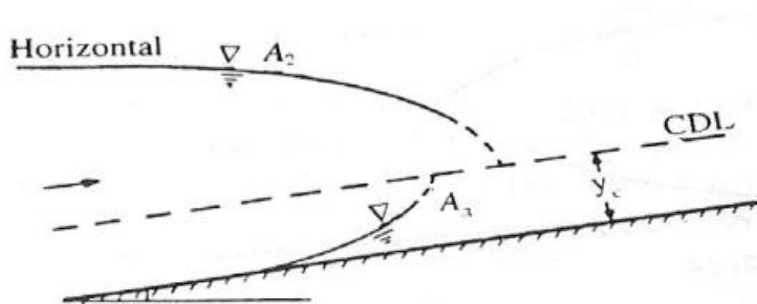
### d). Type H profiles

A horizontal channel can be considered as the lower limit reached by a mild slope as its bed slope becomes flatter. It is obvious that there is no region 1 for a horizontal channel as  $y_0 = \infty$ . The H2 and H3 profiles are similar to M2 and M3 profiles respectively. However, the H2 curve has a horizontal asymptote.



### e). type A profiles

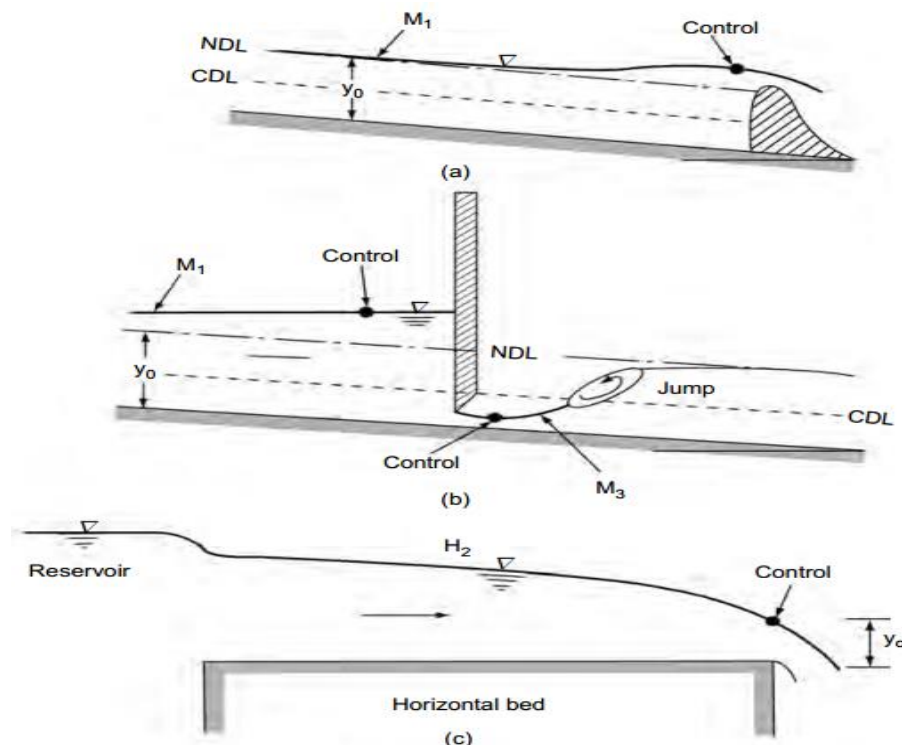
Adverse slopes are rather rare and A2 and A3 curves are similar to H2 and H3 curves respectively. The profiles are of very short length.



## 4.5. Control section

A control section is defined as a section in which a fixed relationship exists between the discharge and depth of flow. Weirs, spillways, sluice gates are some typical examples of structures which give rise to control sections. The critical depth is also a control point. However, it is effective in a flow profile which changes from subcritical to supercritical flow in the reverse case of transition from supercritical flow to subcritical flow; a hydraulic jump is usually formed by passing the critical depth as a control point. Any GVF profile will have at least one control point. This control sections provide a key to the identification of proper profile shapes. A few typical control sections are indicated in the figure below.

*Subcritical flows have controls in the downstream end while supercritical flows are governed by control section existing at the upstream end of the channel section.*



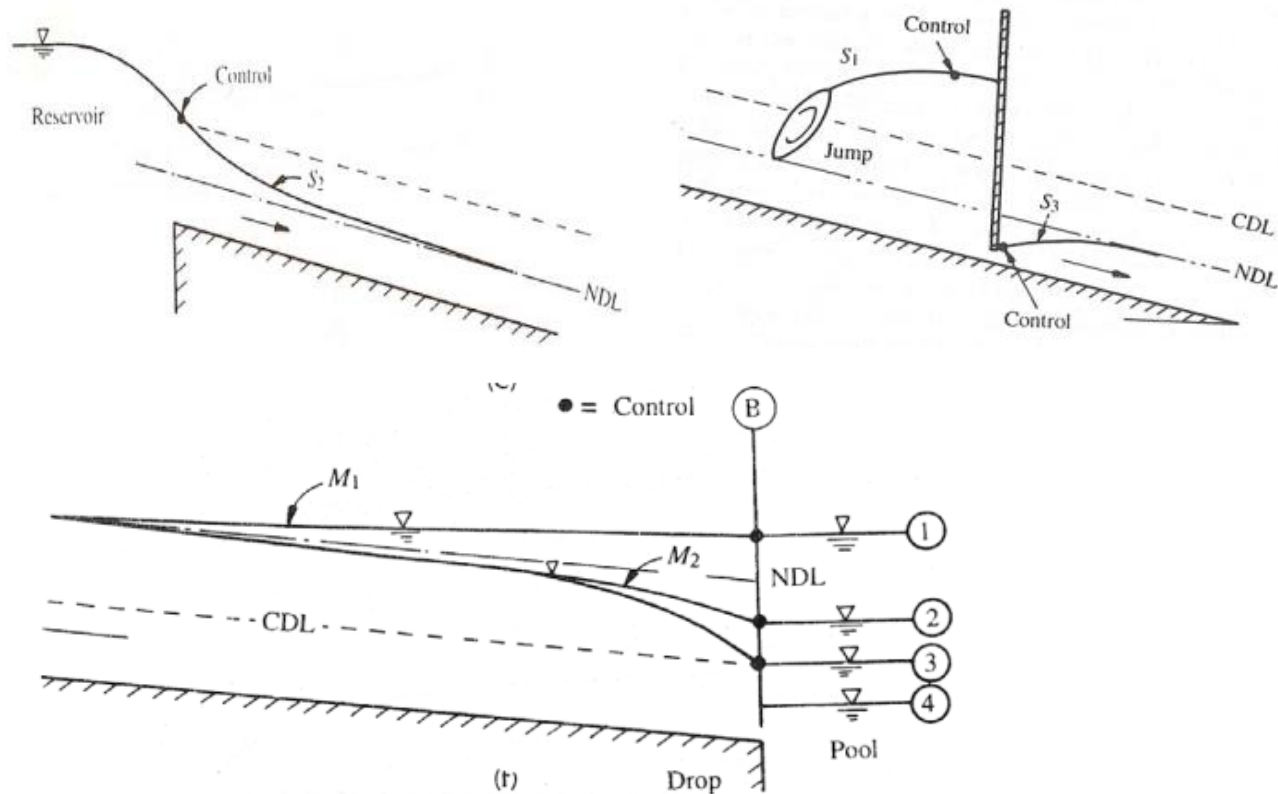


Fig 4.3 examples of control in GVF

#### 4.6. Analysis of flow profiles

The process of identification of possible profiles as a prelude to quantitative computation is known as analysis of flow profiles. A channel carrying a GVF can in general contain different prismatic channel sections of varying hydraulic properties. There can be a number of control sections at a various locations. To determine the resulting water-surface profile in a given case, one should be in a position to analyse the effect of various channel sections and controls connected in series.

#### Break in grade (Serial Combination of Channel Sections)

Simple situation of a serial combination of two channel sections with differing bed slopes are connected in the figure below. The grade changes acts as a control section and this can be classified as a natural control. Various combinations of slopes and the resulting GVF profiles are presented in the figure below. It may be noted that in some situation there can be more than one possible profile.

Procedure to draw the profile of GVF in grade transition:-

1. Draw the longitudinal section of the system.
2. Calculate the critical depth and normal depths of various reaches and draw the CDL and NDL in all reaches. Since  $y_c$  doesn't depend upon the slope CDL will be constant above the channel bed in both slopes.
3. Mark all the controls, both the imposed as well as natural controls.
4. Identify the possible profiles.
5. The normal depth for the mild slope is lower than that of the milder slope in this case the second depth acts as a control.

Various combinations of slopes and the resulting GVF profiles are presented in the figure below. It may be noted that in some situation there can be more than one possible profile.

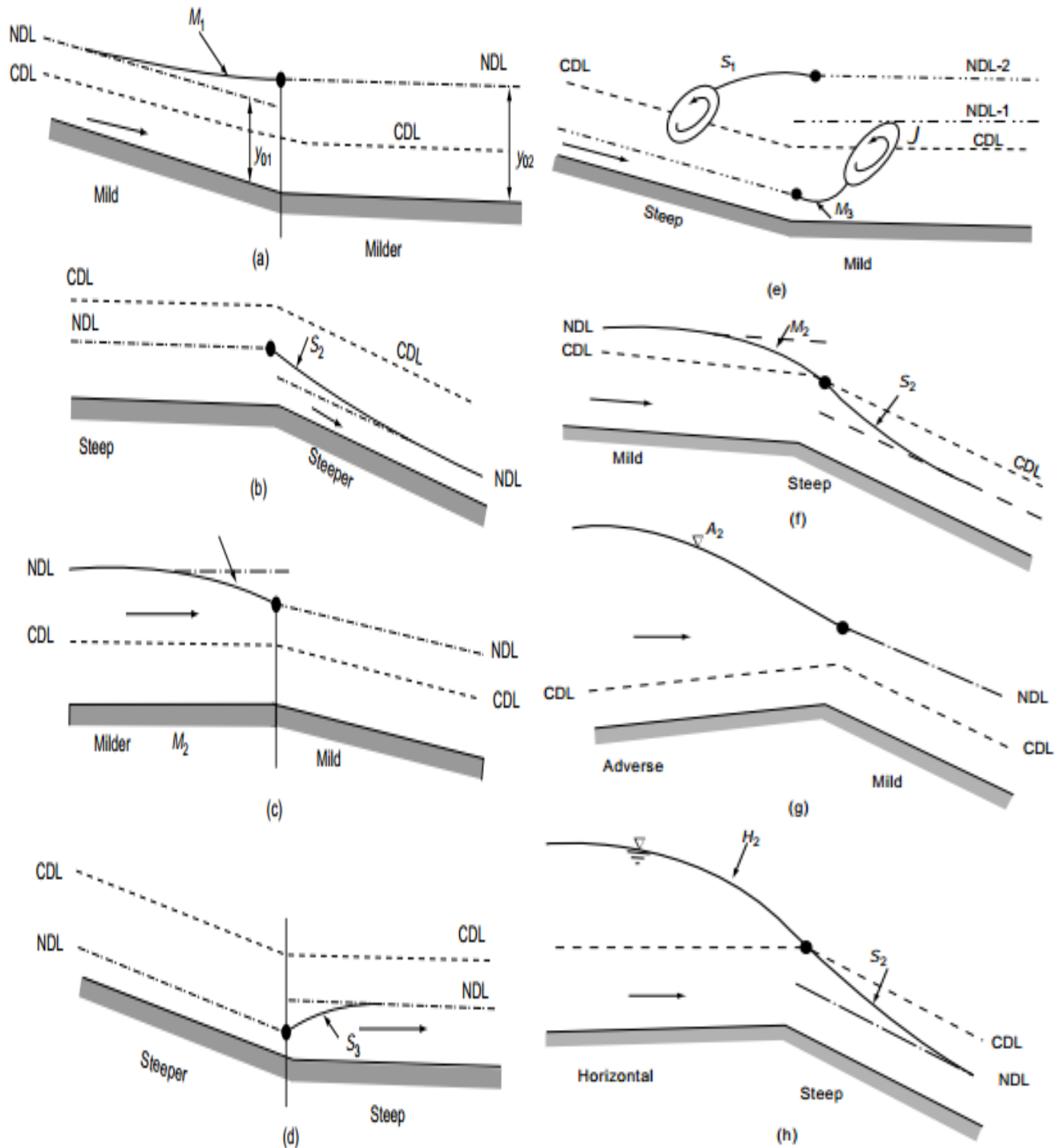


Figure 4.4 GVF profiles at breaking grade

## 4.7 GVF Computations

Almost all major hydraulic engineering activities in the free surface flow involve the computation of GVF profiles. Gradually varied flow involves the solution of the dynamic equation. The main objective is to determine the shape of flow profile. The flow computation is needed to analyze problems such as

- Determination of effect of a hydraulic structure on the channel
- Inundation due to a dam or weir construction
- Estimation of flood zone

Broadly classified there are three Methods of GVF computations

1. Numerical method
2. Direct integration
3. Graphical method

Out of these the graphical method is practically obsolete and is seldom used. Further the numerical method is the most extensively used technique. In the form of a host of available compressive

Software's it is the only method available to solve practical problem in natural channels. The direct integration technique is essentially of academic interest.

### 4.7.1 Numerical method

The numerical solution procedure to solve GVF problems can be broadly classified into two categories as:

#### a. Simple numerical methods

These were developed primarily for hand computation. They usually attempt to solve the energy equation either in the form of the differential energy equation of GVF or in the form of the Bernoulli equation.

#### b. Advanced Numerical Methods

These are normally suitable for use in digital computers as they involve a large number of repeated calculations. They attempt to solve the differential equation of GVF.

Two commonly used simple numerical methods to solve GVF problems are :

1. Direct-step method
2. Standard-step method

#### 4.7.1.1 The Direct Step Method (Distance from Depth)

The direct step method is a simple method applicable to prismatic channels. Depths of flow are specified and the distances between successive depths are calculated. The equation may be used to determine directly (with means explicit) the distance between given differences of depth ( $\Delta y$ ).

Consider the differential energy equation of GVF. —

Writing this in the finite difference form —  $\overline{S_f}$

Where  $\overline{S_f}$  = average friction slope in the reach  $\Delta x$ .

$$\frac{dE}{dx} = \overline{S_0} - \overline{S_f} \dots\dots\dots(4) \text{ and}$$

Between two sections 1 and 2

$$\frac{E_2 - E_1}{\Delta x} = \overline{S_0} - \overline{S_f} \dots\dots\dots(5)$$

This equation is used to calculate the GVF profile.

The hydraulic elements are independent of the distance along the (prismatic) channel. An approximate analysis can be achieved by dividing the channel in a number of successive, short reaches. For each of the reaches the water depth at the beginning can be estimated.

Next the length of reaches can be calculated (step by step) from one end of the reach to the other end. The Chezy or Manning formula is applied to average conditions in each reach to provide an estimate of  $\bar{S}_f$  and  $S_o$ , with the depth and velocity at one end of the reach given, the length can be computed. Depths of flow are specified and the distances between successive depths are calculated.

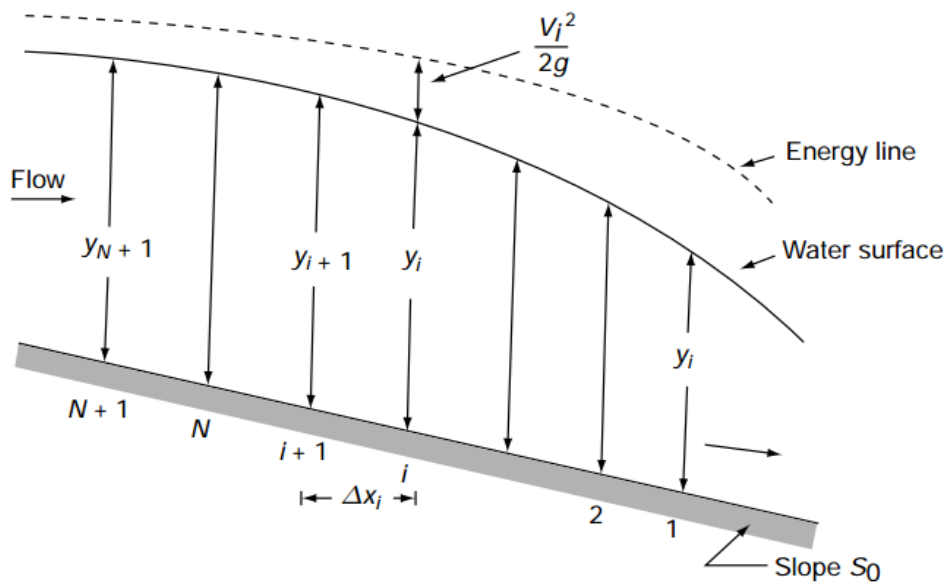


Figure 4.5 the Channel Reach for derivation of direct step method

### Procedure

Referring to figure 4.5, let it be required to find the water surface profile between two section 1 and (N+1) where the depth are  $y_1$  and  $y_{N+1}$  respectively. The channel reach is now divided into N part of known depths, i.e values of  $y_i$   $i=1, N$  are known. It is required to find the distance  $\Delta x_i$  between  $y_i$  and  $y_{i+1}$ . Now between the two sections  $i$  and  $i+1$ .

The sequential evaluation of  $\Delta x_i$  starting from  $i=1$  to  $N$ , will give the distance between the  $N$  sections and thus the GVF profile. The process is explicit and is best done in a tabular manner.

**For the computations the following are needed:**

- ✓ Discharge  $Q$
- ✓ Depth of flow  $y$
- ✓ Area  $A$
- ✓ Hydraulic radius  $R$
- ✓ Roughness coefficient  $n$  or  $C$
- ✓ Coefficient of Coriolis  $\alpha$

1	2	3	4	5	6	7	8	9	10	11	12	13
SL.No	$y(m)$	$A(m^2)$	$P(m)$	$R(m)$	$V(m/s)$	$E(m)$	$\Delta E(m)$	$S_f$	$\bar{S}_f$	$\bar{S}_f$	$\Delta x$	$x$

#### 4.7.1.2 Standard step method

While the direct step method is suitable for use in prismatic channels and there are some basic difficulties in applying it to natural channels. As already indicated in natural channels the cross sectional shapes are likely to vary from section to section and also the cross-section information is known only at a few locations along the channel. Thus the problem of computation of the GVF profile for a natural channel can be stated as:

Given stage at one section it is required to determine the stage at the other section. The sequential determination of the stage as a solution of the above problem will lead to the GVF profile.

The solution of the above problem is obtained by trial and error solution of the basic energy equation. Consider figure 4.6 which shows two sections 1 and 2 at a distance  $\Delta x$ . Calculations are assumed to proceed up stream, equating the total energy at section 1 and 2.

Where  $h_f$  = friction loss and  $h_e$  = eddy loss. The frictional loss  $h_f$  can be estimated as

$$\bar{s}_f$$

Where  $S_f = \frac{V^2}{R^{4/3}}$

There is no rational method for estimating the eddy loss but it is usually expressed as

$$C_e \left| \frac{V_1^2}{2g} - \frac{V_2^2}{2g} \right|$$

Where  $C_e$  is a coefficient having the values as below?.

Nature of transition	Values of coefficient $C_e$	
	Expansion	Contraction
1. No transition (prismatic channel)	0.0	0.0
2. Gradual transition	0.3	0.1
3. Abrupt	0.8	0.6

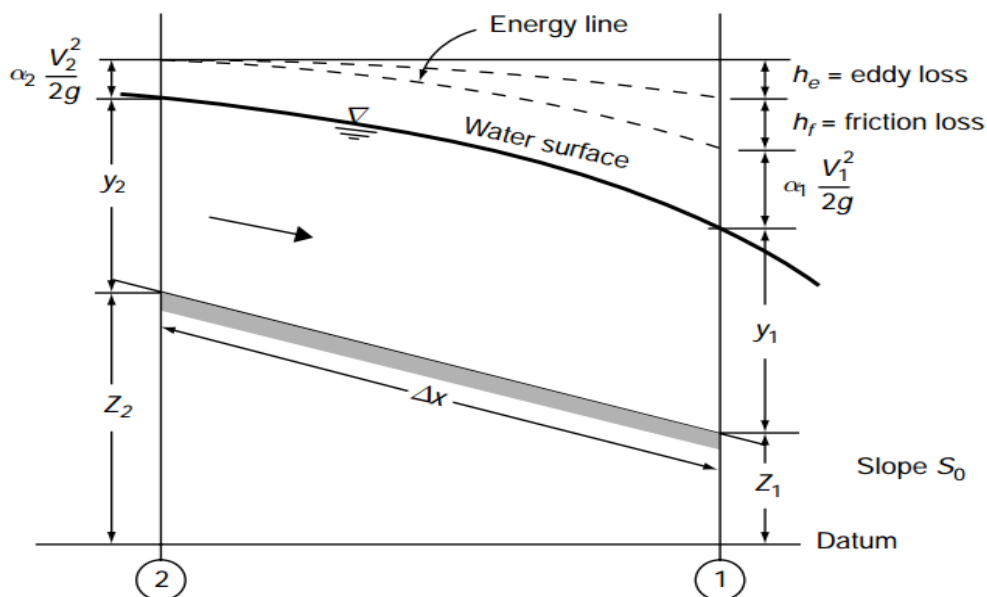


Figure 4.6 the Channel Reach for derivation of Standard step method



Denoting the stage = and the total energy by  $H$ , and using the suffixes 1 and 2 to refer the parameters to appropriate sections,

— and the total energy at section 2 becomes

$$\dots\dots\dots(4.6)$$

The problem can now be stated as: knowing  $H_1$  and the geometry of the channel at sections 1 and it is required to find  $h_2$ . This is achieved in the standard-step method by the trial and error procedure outlined below.

### Procedure

Select a trial value of  $h_2$  and calculating  $H_2$ ,  $h_f$  and  $h_e$  and check whether Eq.(4.6) is satisfied. If there is a difference ,improve the assumed value of  $h_2$  and repeat calculations till the two sides of Eq(4.6) match to an acceptable degree of tolerance.

On the basis of the  $i$ th trial, the  $(i+1)$ th trial value of  $h_2$  can be found by the following procedure suggested by Henderson. Let  $H_E$  be the difference between the left-hand side and right-hand side of Eq(4.6) in the  $i$ th trial,

i.e 
$$\left[ H \quad (H \quad ) \right]$$

The object is to make  $H_E$  vanish by changing the depth  $y_2$ .

Hence 
$$\frac{E}{H} = \frac{y}{H} - \frac{x(S)}{H} \left[ \frac{y}{H} - \frac{y}{H} \right]$$

Since  $y_1$ ,  $Z_1, Z_2$  and  $V_1$  are constants.

$$\frac{y}{H} - \frac{y}{H} \left[ \frac{y}{H} - \frac{y}{H} \right] = \frac{y}{H} - \frac{y}{H} \left[ \frac{y}{H} - \frac{y}{H} \right]$$

Where

For a wide rectangular channel 
$$\frac{y}{H} - \frac{y}{H} \left( \frac{y}{H} \right) = \frac{y}{H} - \frac{y}{H} \left( \frac{y}{H} \right)$$

Hence 
$$\frac{y}{H} - \frac{y}{H} - \frac{y}{H} = \frac{y}{H} - \frac{y}{H} - \frac{y}{H}$$

leading to 
$$\left[ 1 - \frac{y}{H} - \frac{y}{H} \right] = \left[ 1 - \frac{y}{H} - \frac{y}{H} \right]$$

If  $\frac{y}{H} - \frac{y}{H}$  and  $\Delta y_2$  is chosen such that  $\Delta H_E = H_E$

$$\left[ 1 - \frac{y}{H} - \frac{y}{H} \right] = \left[ 1 - \frac{y}{H} - \frac{y}{H} \right]$$

The negative sign denotes that  $\Delta y_2$  is of opposite sign to that of  $H_E$ . it may be noted that if the calculations are performed in the downward direction, as in super critical flow the third term in the denominator will be negative.

For the computation the following data are needed:

- Discharge  $Q$
- Length of the reach  $\Delta x$
- Area  $A$  as function of  $y$
- Hydraulic radius  $R$  as function of  $y$
- Roughness coefficient ( $n$  or  $C$ )
- Carioles coefficient  $\alpha$

The computation of the flow profile by the standard step method is arranged in tabular form

Stat -ion	trial	Elevatio n of bed $z(m)$	$y$ depth ( $m$ )	$h$ or $Z$ stage ( $z+y$ ) ( $m$ )	$A$ ( $m^2$ )	— ( $m$ )	$H$ Total head ( $m$ )	$R$ ( $m$ )	$S_f$ Units of $10^{-4}$	$\bar{S}_f$ Units of $10^{-4}$	Length of reach ( $m$ )	$h_f$ ( $m$ )	$h_e$ ( $m$ )	Total head ( $m$ )	$H_E$ ( $m$ )	( $m$ )
1(a)	1(b)	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16

Each column of the table is explained as follows:

1. The location of the stations is fixed.
2. Water-surface elevation  $Z$  at the station. A trial value is first entered in this column; this will be verified or rejected on the basis of  $h_t$  computations made in the remaining columns of the table. For the first step, this elevation must be given or assumed. In most cases the first entry is known. After this value in the second step has been verified, it becomes the basis for the verification the trial value in the next step, and so on.
3. Depth of flow  $y$  corresponding to the water-surface elevation in col. 2. For instance, the depth of flow  $y$  at the second station is equal to water-surface elevation minus bottom elevation (distance from the first site times bed slope)
4. Water area  $A$  corresponding to  $y$  in col.3
5. Mean velocity  $v$  equal to the given discharge divided by the water area in col. 4
6. Velocity head in  $m$ , corresponding to the velocity col. 5
7. Total head  $E$  computed, equal to the sum of  $Z$  in col. 2 and the velocity head in col. 6
8. Hydraulic radius  $R$  corresponding to  $y$  in col. 3
9. Friction slope  $S_f$  with  $n$  or  $C$ ,  $V$  from col. 5 and  $R$  from col. 8
10. Average friction  $\bar{S}_f$  slope through the reach between the sections in each step, approximately equal to the arithmetic mean of the friction slope just computed in col. 9 and that of the previous step.
11. Length of the reach ( $\Delta x$ ) between the sections.
12. Friction loss in the reach, equal to the product of the values in cols. 10 and 11.
13. transition loss in the reach  $h_e$  (m) —
14. Elevation of the total head  $E$ . this is computed by adding the values of  $h_f$  (and  $h_e$  if calculated in a previous column) in col. 14 to the elevation at the lower end of the reach, which is found in col. 14 of the previous reach. If the value so obtained does not agree closely with that entered in **col. 7**, a new trial value of the water-surface elevation is assumed, and so on, until agreement is obtained. The value that leads to agreement is the correct water-surface elevation. The computation may then proceed to the next step.

### 4.7.2 Direct integration method

The Differential equation of GVF primordial channel given by  $\frac{dy}{dx} = \frac{S_0 - \frac{2}{3} \frac{V^2}{K^2 Z}}{\frac{2}{3} \frac{V^2}{K^2 Z}}$  is a non-linear, first order, ordinary differential equation. Different researchers have made attempts to integrate the GVF equation the method described here but is due to Chow which is based on certain assumptions. Let it be required to find  $y = f(x)$  in the depth range  $y_1$  to  $y_2$  the following two assumptions are made.

1. The conveyance  $K$  at any depth  $y$  is given by  $K^2 = C_2 y^N$  at a normal depth  $y_n$  by  $K_n^2 = C_2 y_n^N$  this implies that in the depth range which includes  $y_1, y_2$  and  $y_n$  the coefficient  $C_2$  and the second hydraulic exponent  $N$  are constants.
2. The section factor  $Z$  at any depth  $y$  is given by  $Z^2 = C_1 y^M$  at a critical depth  $y_c$  by  $Z_c^2 = C_1 y_c^M$  this implies that in the depth range which includes  $y_1, y_2$  and  $y_c$  the coefficient  $C_1$  and the first hydraulic exponent  $M$  are constants.

Substituting the relationship in the differential equation of GVF above gives

$$\frac{dy}{dx} = S_0 \frac{1 - \left(\frac{y_n}{y}\right)^N}{1 - \left(\frac{y_c}{y}\right)^M}$$

Let  $u = \frac{y_n}{y}$  and  $du = -\frac{y_n}{y^2} dy$  then the above equation may be expressed for  $dx$  as

$$dx = \frac{y_n}{S_0} \left[ 1 - \frac{1}{1 - u^N} + \left(\frac{y_c}{y_n}\right)^M \frac{u^{N-M}}{1 - u^N} \right] du$$

$$x = \frac{y_n}{S_0} \left[ u - \int_0^u \frac{du}{1 - u^N} + \left(\frac{y_c}{y_n}\right)^M \int_0^u \frac{u^{N-M}}{1 - u^N} du \right] + \text{const}$$

The first integration on the right side of the above equation is designated by  $F(u, N)$ , or

$$F(u, N) = \int_0^u \frac{du}{1 - u^N} \text{ Is known as the varied flow function.}$$

The second integral in the equation may also be expressed in the form of the varied flow function.

Let  $v = u^{N/J}$  and  $J = N/(N-M+1)$ ; this integral can be transformed into.

$$\int \frac{u^{N-M}}{1 - u^N} du = \frac{J}{N} \int \frac{dv}{1 - v^J} = \frac{N}{J} F(v, J) \quad \text{Where} \quad F(v, J) = \int_0^v \frac{dv}{1 - v^J}$$

Using the notation for varied flow function the equation may be written as

$$x = A \left[ u - F(u, N) + B F(v, J) \right] + \text{const} \quad \text{Where}$$

$$A = \frac{y_n}{S_0}, \quad B = \left(\frac{y_c}{y_n}\right)^M \frac{J}{N}, \quad u = \frac{y}{y_n}, \quad v = u^{N/J} \quad \& \quad J = \frac{N}{N-M+1}$$

And  $F(u, N)$  and  $F(v, J)$  are varied flow functions.

The length of flow profile between two consecutive section 1 and 2 is equal to  $L = x_2 - x_1$

$$L = A \{ (u_2 - u_1) - [F(u_2, N) - F(u_1, N)] + B [F(v_2, J) - F(v_1, J)] \}$$

Where the subscripts 1 and 2 refers to sections 1 and 2, respectively.

The computation of the flow profile by the direct integration method is arranged in tabular form

1	2	3	4	5	6	7	8	9
S.No.	y(m)	u	v	F(u,N)	F(v,J)	x(m)	$\Delta x(m)$	L(m)
1.								

For varied flow function make use of the varied-flow-function table given in tables (at the end of this chapter, appendix D of Ven Te Chow or at the end of chapter 5 of Subramaniya)

#### A. The First Hydraulic Exponent M

In many computations involving a wide range of depth in a channel, such as in the GVF computations, it is convenient on to express the variation of Z with y in an exponential form.

The (Z-y) relationship  $Z^2 = C_1 y^M$  is found to be very advantageous.

Where  $C_1$  = a coefficient and M = an exponent called here as the First hydraulic exponent. it is found that generally M is a slowly – varying function of the aspect ratio for most of the channel shapes.

The value of M for a given channel can be determined by preparing a plot of Z vs y on log-log scale. If M is constant between two points (Z<sub>1</sub>, Y<sub>1</sub>) and (Z<sub>2</sub>, Y<sub>2</sub>) in this plot, it is determined as

$$M = 2 \frac{\log(Z_2/Z_1)}{\log(y_2/y_1)}$$

In the above equation instead of Z a non-dimensionalised Z value can also be used.

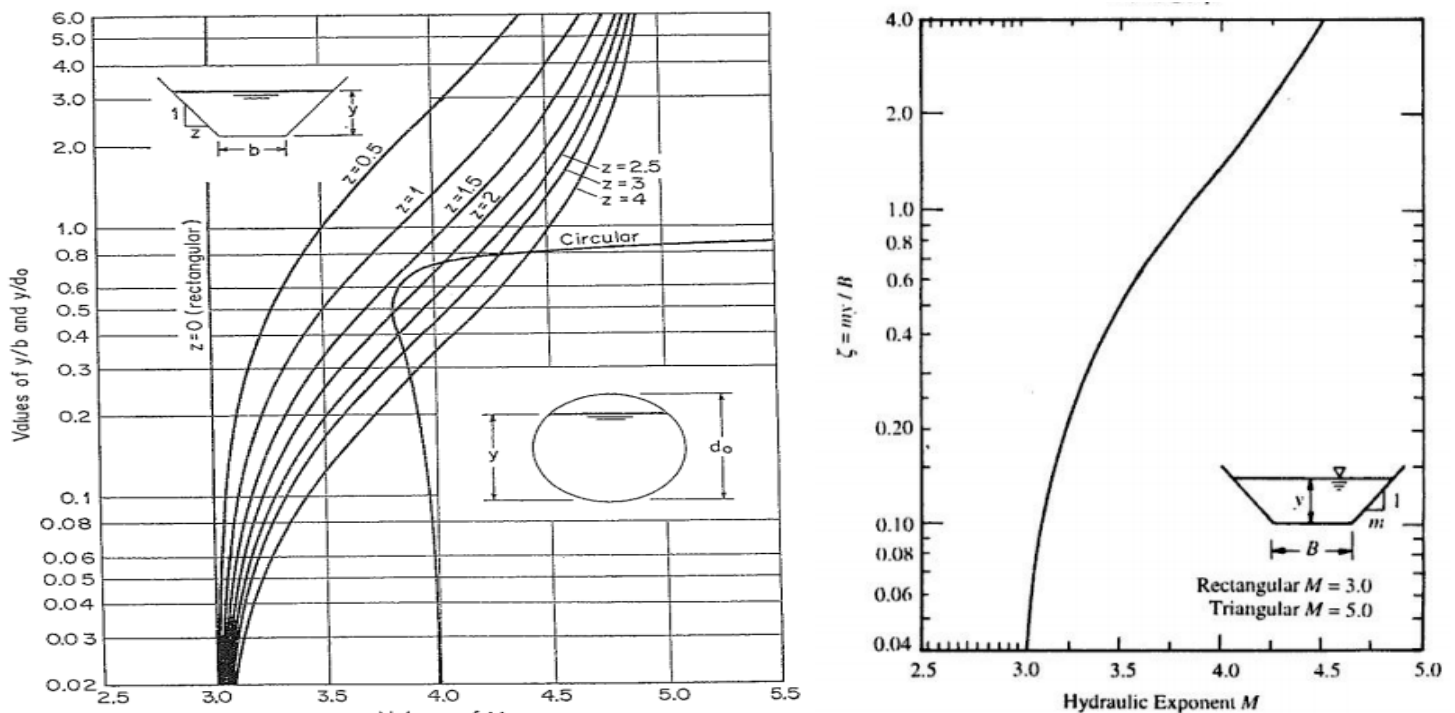


Figure 4.7 variation of first hydraulic exponent M

### B. The Second Hydraulic Exponent $N$

The conveyance of a channel is in general a function of the depth of flow. In calculations involving gradually-varied flow, for purposes of integration, Bakhmeteff introduced the following assumption

$$K^2 = C_2 y^N$$

Where  $C_2$  = a coefficient and  $N$  = an exponent called here as the second hydraulic exponent to distinguish it from the first hydraulic exponent associated with the critical depth. It is found that the second hydraulic exponent  $N$  is essentially constant for a channel over a wide range of depth. Alternatively,  $N$  is usually a slowly varying function of the aspect ratio of the channel.

To determine  $N$  for any channel, a plot of  $\log K$  vs  $\log y$  is prepared. If  $N$  is constant between two point  $(K_1, Y_1)$  and  $(K_2, Y_2)$  in this plot, it is determined as

$$N = 2 \frac{\log(K_1/K_2)}{\log(y_1/y_2)}$$

For a trapezoidal channel, if  $\phi = \frac{AR^{2/3}}{B^{8/3}}$  given in Table 3A.1 is plotted against  $\eta = y/B$  on a log-log paper,

From the slope of the curve at any  $\eta$ , the value of  $N$  at that point can be estimated. Figure 4.7 shows the variation of  $N$  for trapezoidal channels.

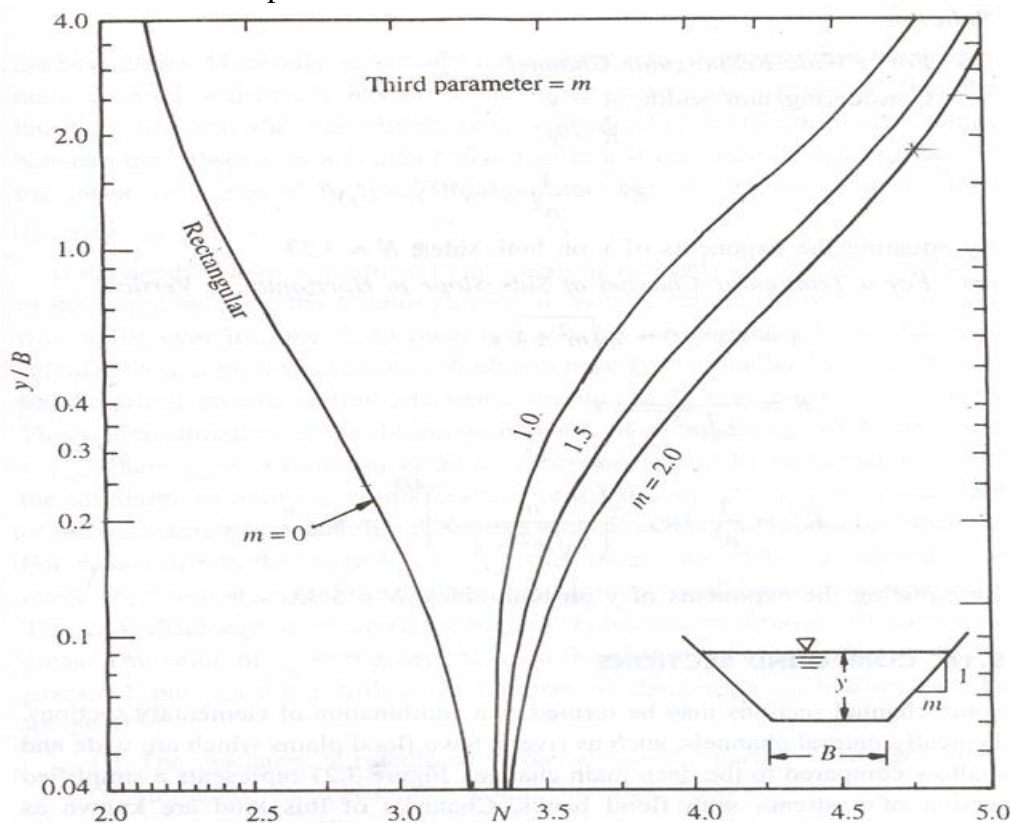


Fig. 4.7 Variation of the second hydraulic exponent  $N$  in trapezoidal channels

The values of  $N$  in this curve have been generated based on the slope of the  $\log K$ - $\log Y$  relation using a computer. Figure 4.7 is useful in the quick estimation of  $N$ . It is seen from this figure that  $N$  is a slowly-varying function of  $y/B$ . For a trapezoidal section, the minimum value of  $N=2.0$  is obtained for a deep rectangular channel and a maximum value of  $N=5.33$  is obtained for a triangular channel.

### 4.7.3 Graphical Integration Method

Consider two channel sections at distance  $x_1$  and  $x_2$  and with corresponding depths of flow  $y_1$  and  $y_2$ . The distance along the channel is  $x$ . If a graph of  $y$  against  $f(y)$  is plotted, then the area under the curve is equivalent to  $x$ . The value of the function  $f(y)$  may be found by substitution of  $A$ ,  $P$ ,  $S_o$  and  $S_f$  for various values of  $y$  and for a given  $Q$ . Hence, the distance  $X$  between the given depths ( $y_1$  and  $y_2$ ) may be calculated (numerical integration) or measured (graphical integration). This numerical/graphical method gives the distance from depth.

This method integrates the equation of gradually varied flow by a numerical procedure.

$$\frac{dy}{dx} = \frac{S_o - S_f}{1 - Fr^2}$$

$$\frac{dx}{dy} = \frac{1 - Fr^2}{S_o - S_f}$$

$$\int_0^x dx = \int_{y_1}^{y_2} \frac{1 - Fr^2}{S_o - S_f} dy$$

$$L = x_2 - x_1 = \int_{y_1}^{y_2} \frac{1 - Fr^2}{S_o - S_f} dy \int_{y_1}^{y_2} \frac{dx}{dy} dy$$

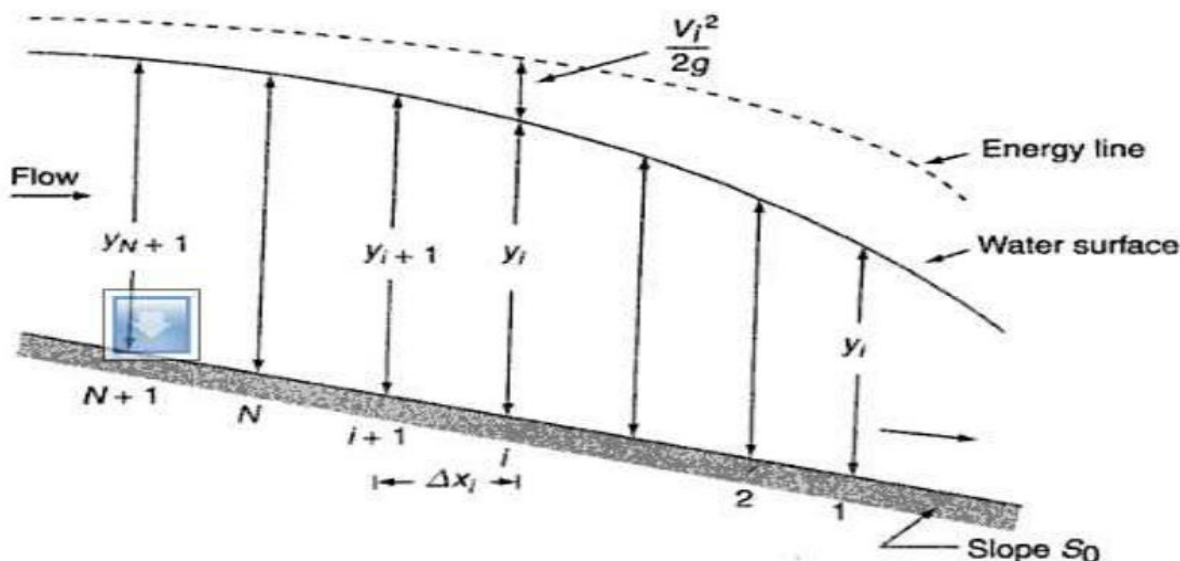


Figure 4.9 The Channel Reach for derivation of Graphical Integration

By this method the largest errors are found in the area with the strongest curvature. This is the region near the control point(s). The accuracy can be improved by varying the steps  $\Delta x$  as a function of the curvature. This method has broad application. It applies to flow in prismatic as well as non-prismatic channels of any shape and slope. The procedure is straightforward and easy to follow. It may become very laborious when applied to actual field problems.



TABLE OF VARIED FLOW FUNCTION ( $F(u,N)$ )

$u N$	2.6	2.8	3.0	3.2	3.4	3.6	3.8	4.0	4.2	4.6	5.0
0.00	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
0.02	0.020	0.020	0.020	0.020	0.020	0.020	0.020	0.020	0.020	0.020	0.020
0.04	0.040	0.040	0.040	0.040	0.040	0.040	0.040	0.040	0.040	0.040	0.040
0.06	0.060	0.060	0.060	0.060	0.060	0.060	0.060	0.060	0.060	0.060	0.060
0.08	0.080	0.080	0.080	0.080	0.080	0.080	0.080	0.080	0.080	0.080	0.080
0.10	0.100	0.100	0.100	0.100	0.100	0.100	0.100	0.100	0.100	0.100	0.100
0.12	0.120	0.120	0.120	0.120	0.120	0.120	0.120	0.120	0.120	0.120	0.120
0.14	0.140	0.140	0.140	0.140	0.140	0.140	0.140	0.140	0.140	0.140	0.140
0.16	0.160	0.160	0.160	0.160	0.160	0.160	0.160	0.160	0.160	0.160	0.160
0.18	0.181	0.180	0.180	0.180	0.180	0.180	0.180	0.180	0.180	0.180	0.180
0.20	0.201	0.200	0.200	0.201	0.200	0.200	0.200	0.200	0.200	0.200	0.200
0.22	0.221	0.221	0.221	0.220	0.220	0.220	0.220	0.220	0.220	0.220	0.220
0.24	0.242	0.241	0.240	0.240	0.240	0.240	0.240	0.240	0.240	0.240	0.240
0.26	0.262	0.262	0.261	0.261	0.261	0.260	0.260	0.260	0.260	0.260	0.260
0.28	0.283	0.282	0.282	0.281	0.281	0.281	0.280	0.280	0.280	0.280	0.280
0.30	0.304	0.303	0.302	0.302	0.301	0.301	0.300	0.300	0.300	0.300	0.300
0.32	0.325	0.324	0.323	0.322	0.322	0.321	0.321	0.321	0.321	0.320	0.320
0.34	0.346	0.344	0.343	0.343	0.342	0.342	0.341	0.341	0.341	0.340	0.340
0.36	0.367	0.366	0.364	0.363	0.363	0.362	0.362	0.361	0.361	0.361	0.360
0.38	0.389	0.387	0.385	0.384	0.383	0.383	0.382	0.382	0.381	0.381	0.381
0.40	0.411	0.408	0.407	0.405	0.404	0.403	0.403	0.402	0.402	0.401	0.401
0.42	0.433	0.430	0.428	0.426	0.425	0.424	0.423	0.423	0.422	0.421	0.421
0.44	0.456	0.452	0.450	0.448	0.446	0.445	0.444	0.443	0.443	0.442	0.441
0.46	0.479	0.475	0.472	0.470	0.468	0.466	0.465	0.464	0.463	0.462	0.462
0.48	0.502	0.497	0.494	0.492	0.489	0.488	0.486	0.485	0.484	0.483	0.482
0.50	0.525	0.521	0.517	0.514	0.511	0.509	0.508	0.506	0.505	0.504	0.503
0.52	0.550	0.544	0.540	0.536	0.534	0.531	0.529	0.528	0.527	0.525	0.523
0.54	0.574	0.568	0.563	0.559	0.556	0.554	0.551	0.550	0.548	0.546	0.544
0.56	0.599	0.593	0.587	0.583	0.579	0.576	0.574	0.572	0.570	0.567	0.565
0.58	0.626	0.618	0.612	0.607	0.603	0.599	0.596	0.594	0.592	0.589	0.587



TABLE OF VARIED FLOW FUNCTION ( $F(u,N)$ ) (continued)

$u/N$	2.6	2.8	3.0	3.2	3.4	3.6	3.8	4.0	4.2	4.6	5.0
0.60	0.653	0.644	0.637	0.631	0.627	0.623	0.620	0.617	0.614	0.611	0.608
0.61	0.667	0.657	0.650	0.644	0.639	0.635	0.631	0.628	0.626	0.622	0.619
0.62	0.680	0.671	0.663	0.657	0.651	0.647	0.643	0.640	0.637	0.633	0.630
0.63	0.694	0.684	0.676	0.669	0.664	0.659	0.655	0.652	0.649	0.644	0.641
0.64	0.709	0.693	0.690	0.683	0.677	0.672	0.667	0.664	0.661	0.656	0.652
0.65	0.724	0.712	0.703	0.696	0.689	0.684	0.680	0.676	0.673	0.667	0.663
0.66	0.738	0.727	0.717	0.709	0.703	0.697	0.692	0.688	0.685	0.679	0.675
0.67	0.754	0.742	0.731	0.723	0.716	0.710	0.705	0.701	0.697	0.691	0.686
0.68	0.769	0.757	0.746	0.737	0.729	0.723	0.718	0.713	0.709	0.703	0.698
0.69	0.785	0.772	0.761	0.751	0.743	0.737	0.731	0.726	0.722	0.715	0.710
0.70	0.802	0.787	0.776	0.766	0.757	0.750	0.744	0.739	0.735	0.727	0.722
0.71	0.819	0.804	0.791	0.781	0.772	0.764	0.758	0.752	0.748	0.740	0.734
0.72	0.836	0.820	0.807	0.796	0.786	0.779	0.772	0.766	0.761	0.752	0.746
0.73	0.854	0.837	0.823	0.811	0.802	0.793	0.786	0.780	0.774	0.765	0.759
0.74	0.868	0.854	0.840	0.827	0.817	0.808	0.800	0.794	0.788	0.779	0.771
0.75	0.890	0.872	0.857	0.844	0.833	0.823	0.815	0.808	0.802	0.792	0.784
0.76	0.909	0.890	0.874	0.861	0.849	0.839	0.830	0.823	0.817	0.806	0.798
0.77	0.930	0.909	0.892	0.878	0.866	0.855	0.846	0.838	0.831	0.820	0.811
0.78	0.950	0.929	0.911	0.896	0.883	0.872	0.862	0.854	0.847	0.834	0.825
0.79	0.971	0.949	0.930	0.914	0.901	0.889	0.879	0.870	0.862	0.849	0.839
0.80	0.994	0.970	0.950	0.934	0.919	0.907	0.896	0.887	0.878	0.865	0.854
0.81	1.017	0.992	0.971	0.954	0.938	0.925	0.914	0.904	0.895	0.881	0.869
0.82	1.041	1.015	0.993	0.974	0.958	0.945	0.932	0.922	0.913	0.897	0.885
0.83	1.067	1.039	1.016	0.996	0.979	0.965	0.952	0.940	0.931	0.914	0.901
0.84	1.094	1.064	1.040	1.019	1.001	0.985	0.972	0.960	0.949	0.932	0.918
0.85	1.121	1.091	1.065	1.043	1.024	1.007	0.993	0.980	0.969	0.950	0.935
0.86	1.153	1.119	1.092	1.068	1.048	1.031	1.015	1.002	0.990	0.970	0.954
0.87	1.182	1.149	1.120	1.095	1.074	1.055	1.039	1.025	1.012	0.990	0.973
0.88	1.228	1.181	1.151	1.124	1.101	1.081	1.064	1.049	1.035	1.012	0.994
0.89	1.255	1.216	1.183	1.155	1.131	1.110	1.091	1.075	1.060	1.035	1.015

TABLE OF VARIED FLOW FUNCTION ( $F(u,N)$ ) continued

$u N$	2.6	2.8	3.0	3.2	3.4	3.6	3.8	4.0	4.2	4.6	5.0
0.90	1.294	1.253	1.218	1.189	1.163	1.140	1.120	1.103	1.087	1.060	1.039
0.91	1.338	1.294	1.257	1.225	1.197	1.173	1.152	1.133	1.116	1.088	1.064
0.92	1.351	1.340	1.300	1.266	1.236	1.210	1.187	1.166	1.148	1.117	1.092
0.93	1.435	1.391	1.348	1.311	1.279	1.251	1.226	1.204	1.184	1.151	1.123
0.94	1.504	1.449	1.403	1.363	1.328	1.297	1.270	1.246	1.225	1.188	1.158
0.950	1.582	1.518	1.467	1.423	1.385	1.352	1.322	1.296	1.272	1.232	1.199
0.960	1.665	1.601	1.545	1.497	1.454	1.417	1.385	1.355	1.329	1.285	1.248
0.970	1.780	1.707	1.644	1.590	1.543	1.501	1.464	1.431	1.402	1.351	1.310
0.975	1.853	1.773	1.707	1.649	1.598	1.554	1.514	1.479	1.447	1.393	1.348
0.980	1.946	1.855	1.783	1.720	1.666	1.617	1.575	1.536	1.502	1.443	1.395
0.985	2.056	1.959	1.880	1.812	1.752	1.699	1.652	1.610	1.573	1.508	1.454
0.990	2.212	2.106	2.017	1.940	1.973	1.814	1.761	1.714	1.671	1.598	1.537
0.995	2.478	2.355	2.250	2.159	2.079	2.008	1.945	1.889	1.838	1.751	1.678
0.999	3.097	2.931	2.788	2.663	2.554	2.457	2.370	2.293	2.223	2.102	2.002
1.000	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$	$\infty$
1.001	2.640	2.399	2.184	2.008	1.856	1.725	1.610	1.508	1.417	1.264	1.138
1.005	2.022	1.818	1.649	1.506	1.384	1.279	1.188	1.107	1.036	0.915	0.817
1.010	1.757	1.572	1.419	1.291	1.182	1.089	1.007	0.936	0.873	0.766	0.681
1.015	1.602	1.428	1.286	1.166	1.065	0.978	0.902	0.836	0.778	0.680	0.602
1.020	1.493	1.327	1.191	1.078	0.982	0.900	0.828	0.766	0.711	0.620	0.546
1.03	1.340	1.186	1.060	0.955	0.866	0.790	0.725	0.668	0.618	0.535	0.469
1.04	1.232	1.086	0.967	0.868	0.785	0.714	0.653	0.600	0.554	0.477	0.415
1.05	1.150	1.010	0.896	0.802	0.723	0.656	0.598	0.548	0.504	0.432	0.374
1.06	1.082	0.948	0.838	0.748	0.672	0.608	0.553	0.506	0.464	0.396	0.342
1.07	1.026	0.896	0.790	0.703	0.630	0.569	0.516	0.471	0.431	0.366	0.315
1.08	0.978	0.851	0.749	0.665	0.595	0.535	0.485	0.441	0.403	0.341	0.292
1.09	0.935	0.812	0.713	0.631	0.563	0.506	0.457	0.415	0.379	0.319	0.272
1.10	0.897	0.777	0.681	0.601	0.536	0.480	0.433	0.392	0.357	0.299	0.254
1.11	0.864	0.746	0.652	0.575	0.511	0.457	0.411	0.372	0.338	0.282	0.239
1.12	0.883	0.718	0.626	0.551	0.488	0.436	0.392	0.354	0.321	0.267	0.225

TABLE OF VARIED FLOW FUNCTION ( $F(u,N)$ ) continued

$u/N$	2.6	2.8	3.0	3.2	3.4	3.6	3.8	4.0	4.2	4.6	5.0
1.13	0.805	0.692	0.602	0.529	0.468	0.417	0.374	0.337	0.305	0.253	0.212
1.14	0.780	0.669	0.581	0.509	0.450	0.400	0.358	0.322	0.291	0.240	0.201
1.15	0.756	0.647	0.561	0.490	0.432	0.384	0.343	0.308	0.278	0.229	0.291
1.16	0.734	0.627	0.542	0.473	0.417	0.369	0.329	0.295	0.266	0.218	0.181
1.17	0.713	0.608	0.525	0.458	0.402	0.356	0.317	0.283	0.255	0.208	0.173
1.18	0.694	0.591	0.509	0.443	0.388	0.343	0.305	0.272	0.244	0.199	0.165
1.19	0.676	0.574	0.494	0.429	0.375	0.331	0.294	0.262	0.235	0.191	0.157
1.20	0.659	0.559	0.480	0.416	0.363	0.320	0.283	0.252	0.226	0.183	0.150
1.22	0.628	0.531	0.454	0.392	0.341	0.299	0.264	0.235	0.209	0.168	0.138
1.24	0.600	0.505	0.431	0.371	0.322	0.281	0.248	0.219	0.195	0.156	0.127
1.26	0.574	0.482	0.410	0.351	0.304	0.265	0.233	0.205	0.182	0.145	0.117
1.28	0.551	0.461	0.391	0.334	0.288	0.250	0.219	0.193	0.170	0.135	0.108
1.30	0.530	0.442	0.373	0.318	0.274	0.237	0.207	0.181	0.160	0.126	0.100
1.32	0.510	0.424	0.357	0.304	0.260	0.225	0.196	0.171	0.150	0.118	0.093
1.34	0.492	0.408	0.342	0.290	0.248	0.214	0.185	0.162	0.142	0.110	0.087
1.36	0.475	0.393	0.329	0.278	0.237	0.204	0.176	0.153	0.134	0.103	0.081
1.38	0.459	0.378	0.316	0.266	0.226	0.194	0.167	0.145	0.127	0.097	0.076
1.40	0.444	0.365	0.304	0.256	0.217	0.185	0.159	0.138	0.120	0.092	0.071
1.42	0.431	0.353	0.293	0.246	0.208	0.177	0.152	0.131	0.114	0.087	0.067
1.44	0.417	0.341	0.282	0.236	0.199	0.169	0.145	0.125	0.108	0.082	0.063
1.46	0.405	0.330	0.273	0.227	0.191	0.162	0.139	0.119	0.103	0.077	0.059
1.48	0.394	0.320	0.263	0.219	0.184	0.156	0.133	0.113	0.098	0.073	0.056
1.50	0.383	0.310	0.255	0.211	0.177	0.149	0.127	0.108	0.093	0.069	0.053
1.55	0.358	0.288	0.235	0.194	0.161	0.135	0.114	0.097	0.083	0.061	0.046
1.60	0.335	0.269	0.218	0.179	0.148	0.123	0.103	0.087	0.074	0.054	0.040



TABLE OF VARIED FLOW FUNCTION ( $F(u,N)$ ) continued

$u/N$	2.6	2.8	3.0	3.2	3.4	3.6	3.8	4.0	4.2	4.6	5.0
1.65	0.316	0.251	0.203	0.165	0.136	0.113	0.094	0.079	0.067	0.048	0.035
1.70	0.298	0.236	0.189	0.153	0.125	0.103	0.086	0.072	0.060	0.043	0.031
1.75	0.282	0.222	0.177	0.143	0.116	0.095	0.079	0.065	0.054	0.038	0.027
1.80	0.267	0.209	0.166	0.133	0.108	0.088	0.072	0.060	0.049	0.034	0.024
1.85	0.254	0.198	0.156	0.125	0.100	0.082	0.067	0.055	0.045	0.031	0.022
1.90	0.242	0.188	0.147	0.117	0.094	0.076	0.062	0.050	0.041	0.028	0.020
1.95	0.231	0.178	0.139	0.110	0.088	0.070	0.057	0.046	0.038	0.026	0.018
2.00	0.221	0.169	0.132	0.104	0.082	0.066	0.053	0.043	0.035	0.023	0.016
2.10	0.202	0.154	0.119	0.092	0.073	0.058	0.046	0.037	0.030	0.019	0.013
2.20	0.186	0.141	0.107	0.083	0.065	0.051	0.040	0.032	0.025	0.016	0.011
2.3	0.173	0.129	0.098	0.075	0.058	0.045	0.035	0.028	0.022	0.014	0.009
2.4	0.160	0.119	0.089	0.068	0.052	0.040	0.031	0.024	0.019	0.012	0.008
2.5	0.150	0.110	0.082	0.062	0.047	0.036	0.028	0.022	0.017	0.010	0.006
2.6	0.140	0.102	0.076	0.057	0.043	0.033	0.025	0.019	0.015	0.009	0.005
2.7	0.131	0.095	0.070	0.052	0.039	0.029	0.022	0.017	0.013	0.008	0.005
2.8	0.124	0.089	0.065	0.048	0.036	0.027	0.020	0.015	0.012	0.007	0.004
2.9	0.117	0.083	0.060	0.044	0.033	0.024	0.018	0.014	0.010	0.006	0.004
3.0	0.110	0.078	0.056	0.041	0.030	0.022	0.017	0.012	0.009	0.005	0.003
3.5	0.085	0.059	0.041	0.029	0.021	0.015	0.011	0.008	0.006	0.003	0.002
4.0	0.069	0.046	0.031	0.022	0.015	0.010	0.007	0.005	0.004	0.002	0.001
4.5	0.057	0.037	0.025	0.017	0.011	0.008	0.005	0.004	0.003	0.001	0.001
5.0	0.048	0.031	0.020	0.013	0.009	0.006	0.004	0.003	0.002	0.001	0.000
6.0	0.036	0.022	0.014	0.009	0.006	0.004	0.002	0.002	0.001	0.000	0.000
7.0	0.028	0.017	0.010	0.006	0.004	0.002	0.002	0.001	0.001	0.000	0.000
8.0	0.022	0.013	0.008	0.005	0.003	0.002	0.001	0.001	0.000	0.000	0.000
9.0	0.019	0.011	0.006	0.004	0.002	0.001	0.001	0.000	0.000	0.000	0.000
10.0	0.016	0.009	0.005	0.003	0.002	0.001	0.001	0.000	0.000	0.000	0.000
20.0	0.011	0.006	0.002	0.001	0.001	0.000	0.000	0.000	0.000	0.000	0.000